Observations of the Diurnal Cycle of Marine Stratocumulus During the VOCALS Regional Experiment

NC STATE UNIVERSITY

Clouds and Precipitation Processes and Patterns Group

Casey D. Burleyson¹, Sandra E. Yuter¹, and Simon P. de Szoeke²

¹ Department of Marine, Earth, and Atmospheric Sciences, North Carolina State University ² College of Atmospheric and Oceanic Sciences, Oregon State University

I) Motivation

The cloud-topped boundary lavers (CTBL) in the SE Pacific exhibit a strong diurnal cycle because of vertical mixing driven by cloud-top radiational cooling. The diurnal evolution of the cloud deck has been previously documented in numerous observational and modeling studies. The VOCALS Regional Experiment obtained a unique ship-based data set in which multiple facets of the diurnal cycle were measured simultaneously. Our work documents new details of diurnal variations in marine stratocumulus based on data sets collected on the NOAA ship Ronald H. Brown (RHB).

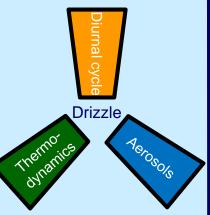


Figure 1 - The diurnal cycle, aerosols, and thermodynamics are three key forcings that jointly determine the characteristics of the CTBL system. Detailed documentation of the diurnal cycle and its variation with longitude will aid in isolating the roles of aerosols and thermodynamics

II) Instruments

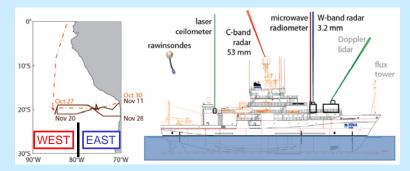


Figure 2 - Data from multiple legs across the 20°S parallel are utilized and divided into two subsets based on longitude (W and E of 80°W). Instrumentation on the RHB that are of primary importance to this study are: the vertically-pointing W-band and scanning C-band radars, the laser ceilometer, flux tower, and rawinsondes which were launched nominally every four hours during the cruise.

III) Surface Observations

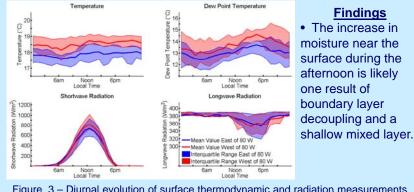


Figure 3 – Diurnal evolution of surface thermodynamic and radiation measurements divided into eastern and western sides of the domain (see Fig. 2). Solid lines indicate the hourly mean value and shaded regions outline the area between the 25th and 75th nercentiles

IV) Cloud Observations

Findings

Cloud top height decreases

throughout the afternoon and

Cloud thickness decreases

significantly after sunrise and

middle of the afternoon.

layer in the west.

radar reflectivity over more than a 9000 km² area.

Underpinning

Findings

Cloud thickness is similar

both east and west of 80°W

despite a deeper boundary

Distribution East of 80° W Distribution West of 80° W

begins to increase again in the

begins to increase near sunset.

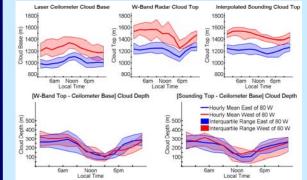
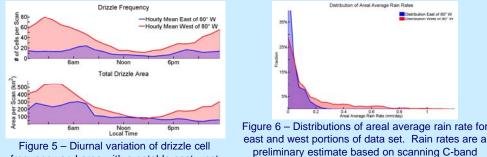


Figure 4 – Hourly average values for stratocumulus clouds over the ship of the cloud base height, top height and thickness.

V) Precipitation Observations



frequency and area with a notable east-west division

Findings

 Drizzle increases through the late afternoon into the evening, maximizes near 3 am, and decreases sharply with sunrise. The decrease is earlier and slower in the west compared to the east.

 Daytime minimum observations of drizzle frequency and area are approximately 20% of the overnight maximum values.

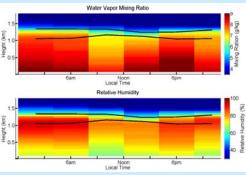
 Heavy drizzle occurs more frequently in the western portion of the domain compared to the eastern portion.

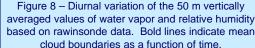
VI) Boundary Layer Coupling Observations

Cloud Base - LCL Height) East of 80° W # of 10 Minute Samples The difference between the cloud base and LCL of a lifted surface parcel will be smallest when the boundary layer is most coupled. Noon ocal Time Noon ocal Time (Cloud Base - I Cl. Height) Distributions I CL Height\ C ervations · The boundary layer is less Single 10 Minute Sample Distribution East of 80° W coupled during the day and quickly couples shortly after sunset Larger differences between cloud base and LCL heights are Figure 7 – Diurnal and longitudinal variation of a coupling more frequent in the western portion of the domain.

measurement based on lifted surface parcels.

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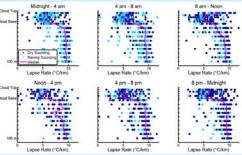
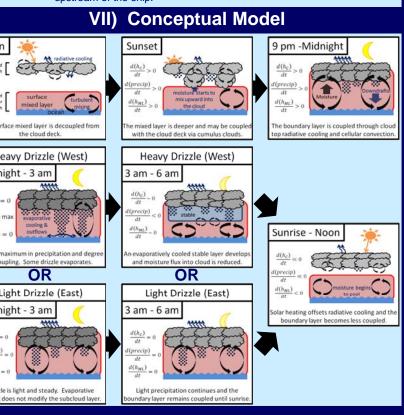
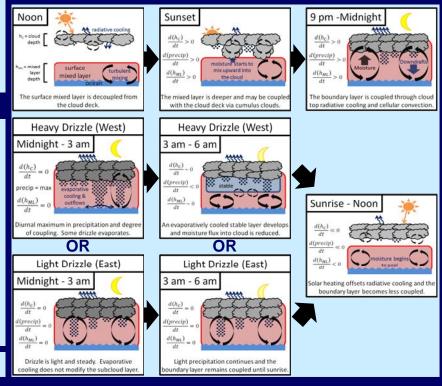
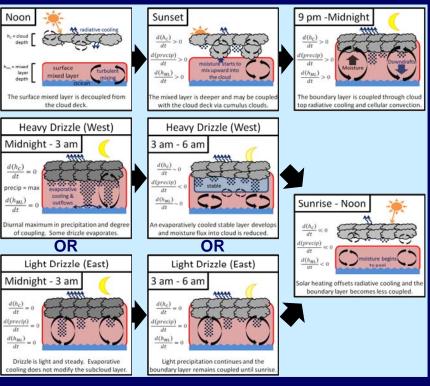


Figure 9 – Measurements of the cloud relative lapse rates. Dark blue dots indicate soundings that were likely taken in dry conditions whereas light blue observations were likely taken with significant drizzle upstream of the ship.



he surface mixed layer is decoupled from







VI) Boundary Layer Coupling (continued)

cloud boundaries as a function of time.

Findings

 Water vapor increases in the lower half of the boundary layer during the afternoon. Stronger boundary layer coupling in the evening allows some of this moisture to mix from the surface into the cloud. lowering cloud base.

 Evaporation of drizzle within downdrafts adds based on rawinsonde data. Bold lines indicate mean moisture to the lower portion of the boundary laver.

> Smaller subcloud lapse rates overnight may be evidence for the modification of the boundary layer by drizzle.

> · An evaporatively cooled stable subcloud layer can develop overnight during precipitation. Such a stable laver would reduce moisture flux into the cloud and reduce precipitation (Bretherton and Wvant 1997).